

Observationally Closing the Gap Between IR Radiative  
Forcing and Changes in IR Radiation Climate  
or  
Is Atmospheric Infrared Radiation Doing What is Supposed to Do?

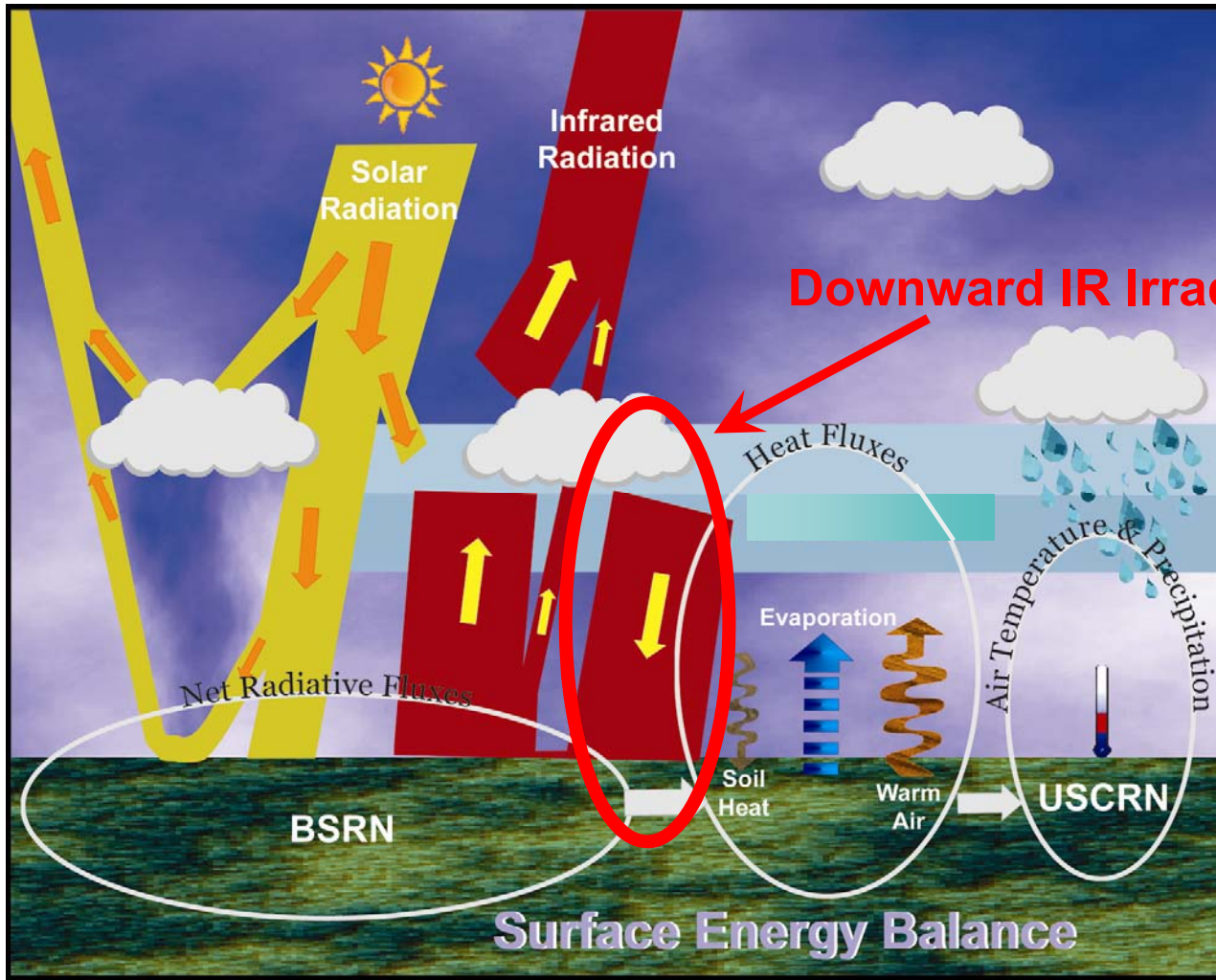
Ellsworth G. Dutton and the ESRL/GMD Radiation Group  
NOAA, ESRL Boulder, Colorado 80305

With thanks to:

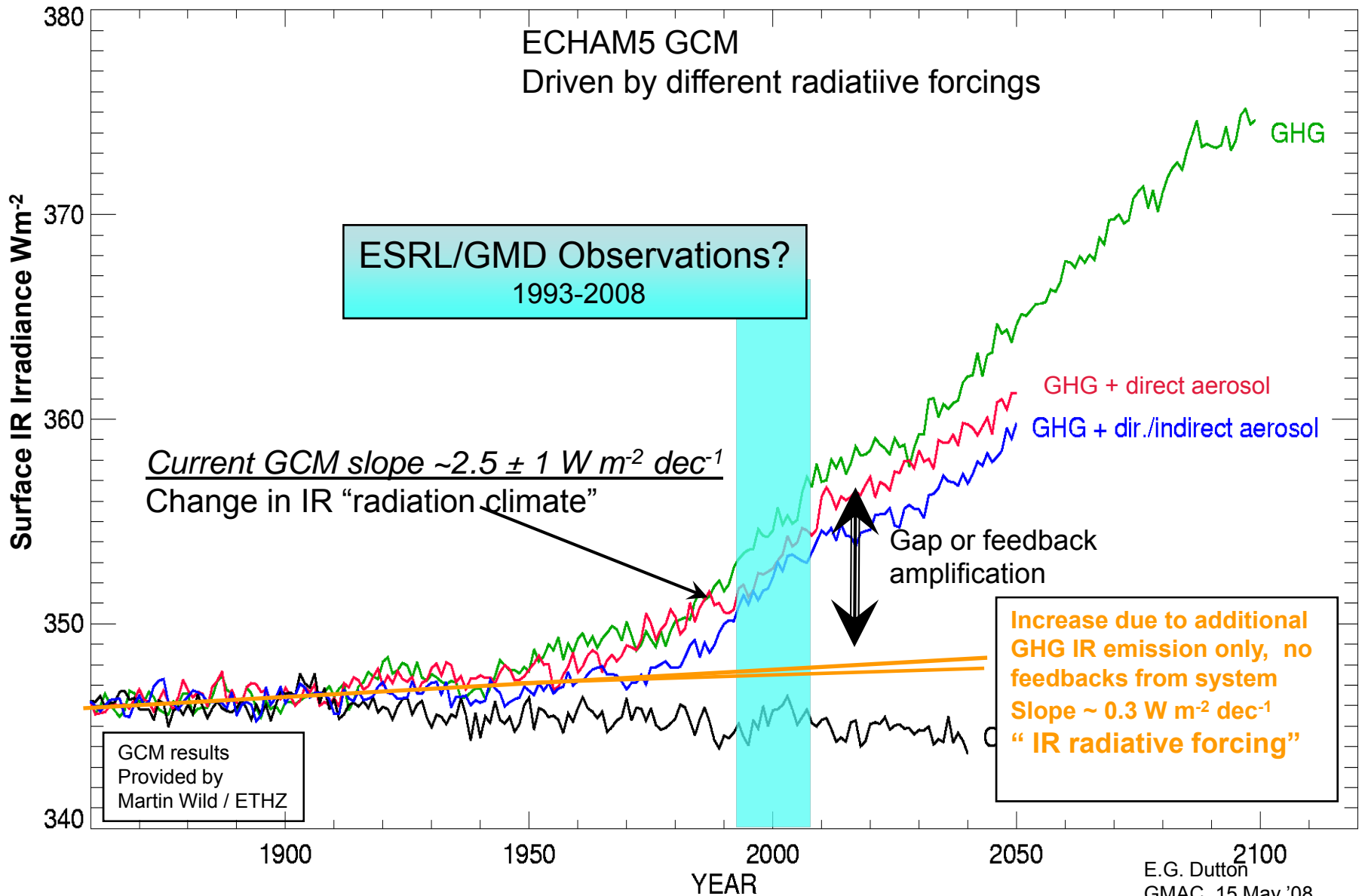
Martin Wild (ETHZ- ECHAM), Norm Wood (CSU- NCAR/CCSM, B Collins),  
Stuart Freidenreich (GFDL-CM2, Delworth) for GCM results

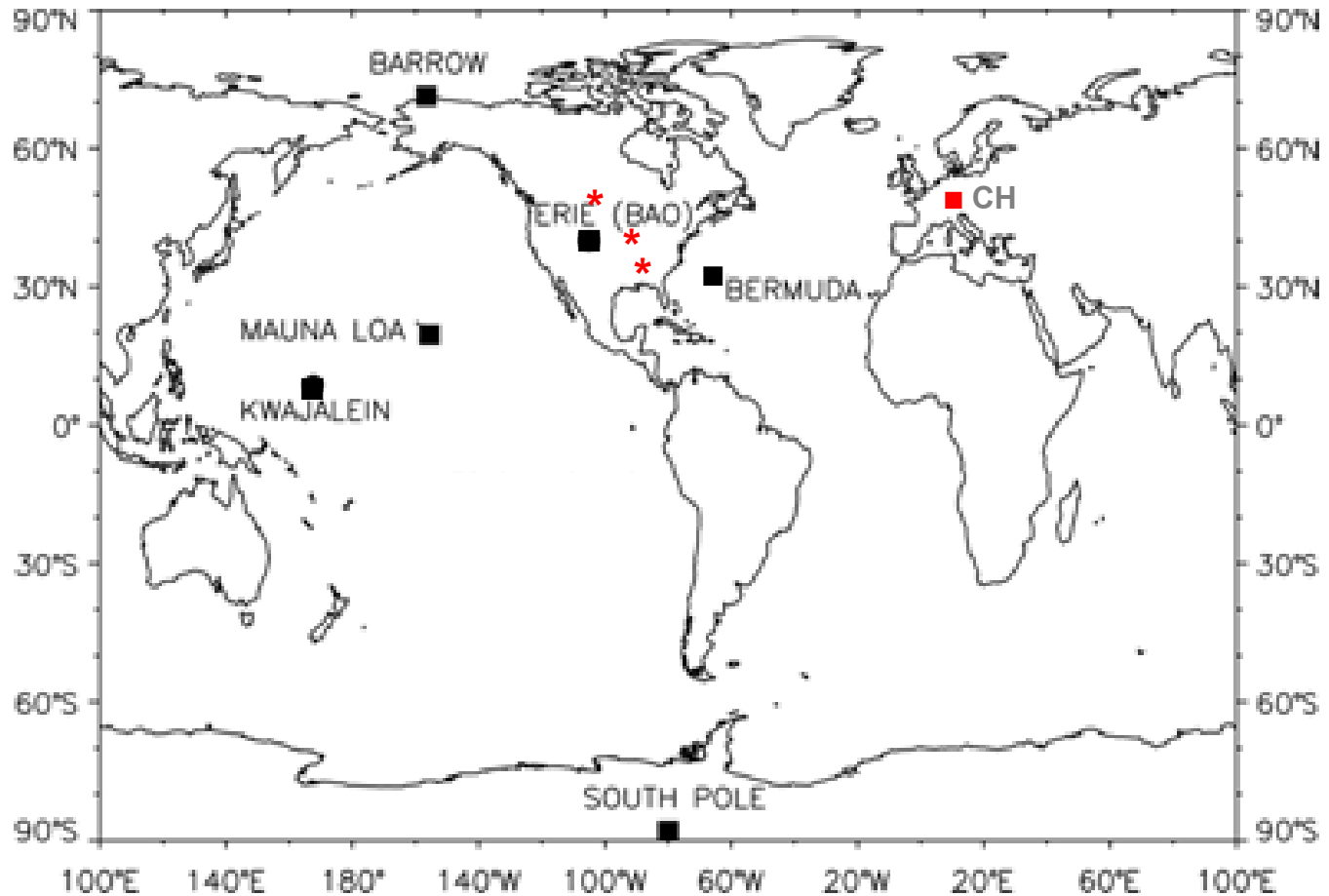
# Terminology:

- Downward IR (LW) Irradiance at the Earth's Surface is quantity of interest in this talk
- Downward IR (LW) Irradiance at the Earth's Surface is the integrated radiant power emitted downward by the atmosphere between about  $3.5 \mu\text{m}$  –  $100 \mu\text{m}$  and intercepted on a horizontal plane at the Earth's surface. It is the combined natural and anthropogenic “greenhouse” radiation, f(T, GHG, H<sub>2</sub>O, Clds, aerosols) global annual mean  $\sim 350 \text{ W m}^{-2}$
- Longwave (LW), infrared (IR), Terrestrial IR, Thermal IR, IR irradiance, and IR radiation may be used interchangeably in this talk
- IR anomalies – Deseasonalized with long-term mean subtracted.



# Global Mean Downwelling Longwave Radiation at the Earth's Surface



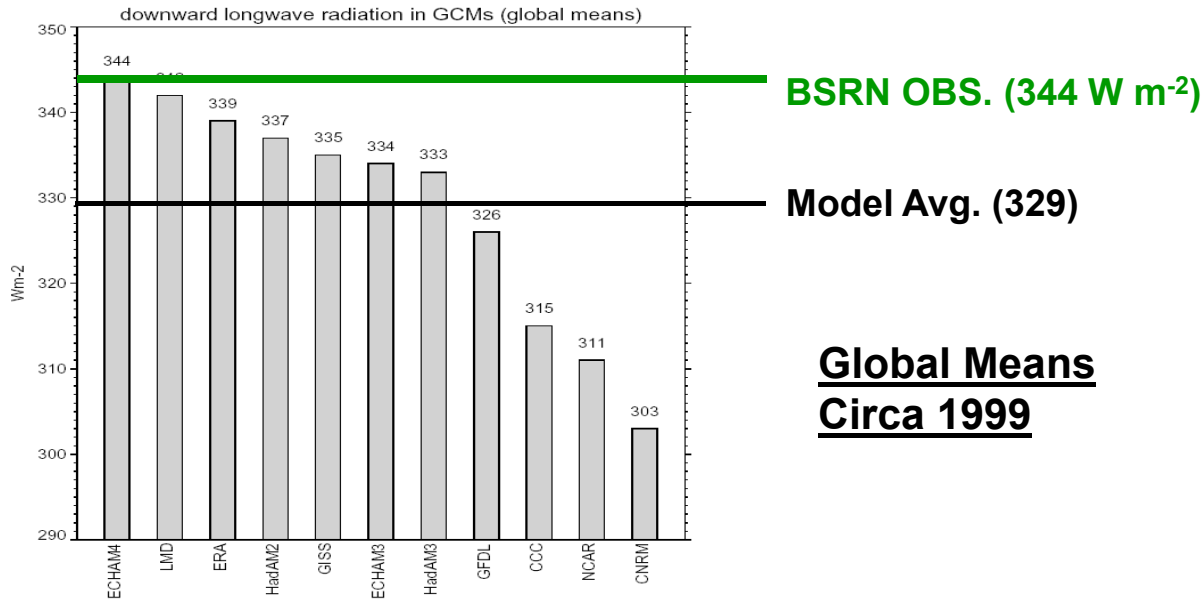


- ESRL-GMD Global Baseline (1993-2008)
- \* ESRL-GMD SURFRAD (1995 – 2008)
- Swiss network (1995-2002, R. Philipona et al. 2005)

# ESRL-GMD Surface IR Observations: A few details (G-Rad global baseline network, 1993 - 2008)

- Commercial **pyrgeometers**
- **Albrecht & Cox** calibration and data reduction methodology
- Calibration **accuracy**  $\sim 3 \text{ W m}^{-2}$ , **traceable int'l**
- Calibration **stability**  $< 0.2\%$  ( $0.7 \text{ W m}^{-2}$ )  $\text{dec}^{-1}$
- Field calibration frequency **once per 1 – 3 years**
- **Continuous** sampling
- Manually edited and reviewed
- Subsequent **analyses**:
  - Deseasonalized 1-day averages  $\rightarrow$  20-day averages  $\rightarrow$  AR-1 residuals
  - Two trend or analyses then applied:
    - **Linear regression**
    - **Mann-Kendall tests on Sens slopes**
  - **Variance reduction** from combining remote sample sites

**GCM surface IR agreement with observations**  
**M. Wild et al., 2001 (see Wild et al 2005 for update)**



E.G. Dutton  
 GMAC, 15 May '08  
 Boulder, Colo.

**BEFORE**

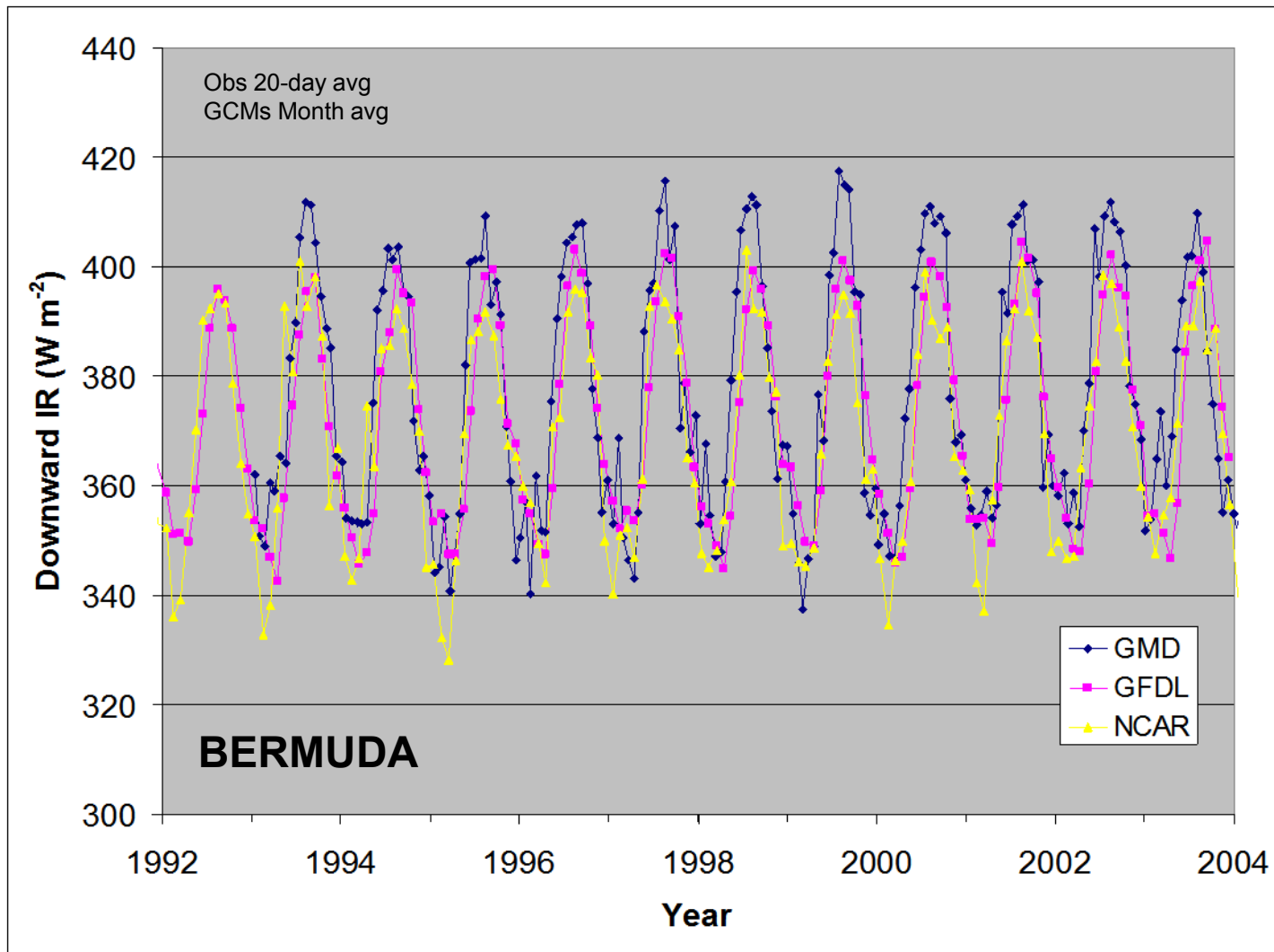
**AFTER**

**GCM grid cell & GMD Obs averages 1993 – 2004 (W m<sup>-2</sup>)**

	Barrow	Boulder (Erie)	Bermuda	Mauna Loa	Kwaj.	S. Pole	Global
CCSM	249.5	266.2	369.3	386.1	420.8	108.0	340
GFDL	243.5	289.1	372.1	390.3	420.9	107.2	338
ECHAM4	238.0	294.4	392.0	-	440.0	113.8	344
<b>OBS</b>	<b>238.3</b>	<b>291.7</b>	<b>377.1</b>	<b>236.4</b>	<b>421.4</b>	<b>111.7</b>	<b>344</b>

Within  
 ~5 W m<sup>-2</sup>  
 of Obs

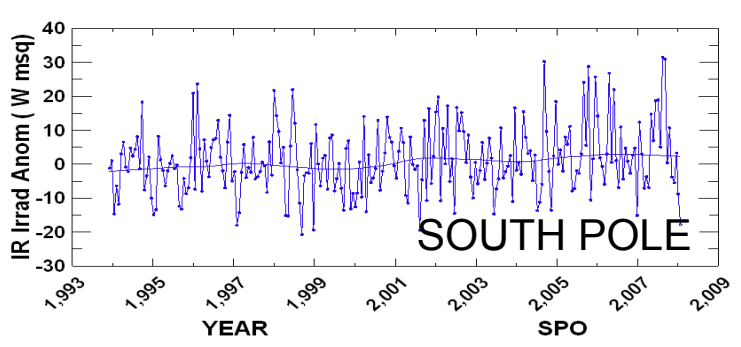
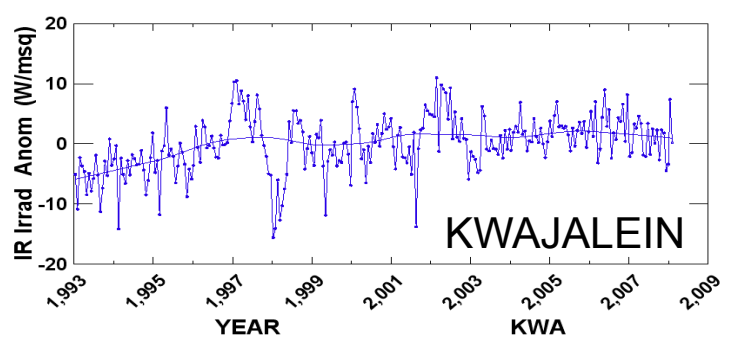
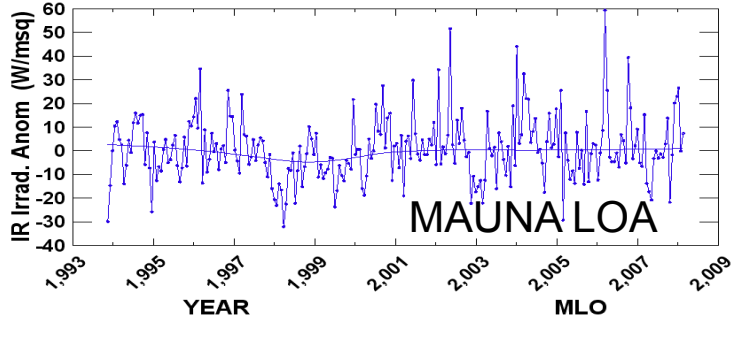
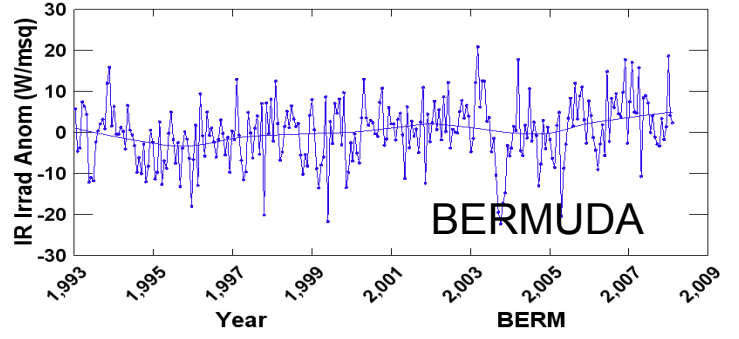
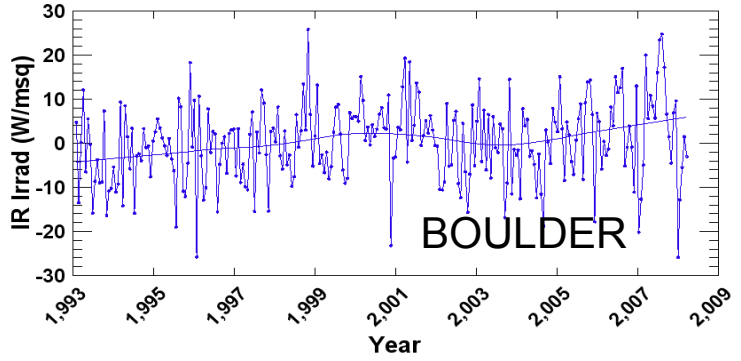
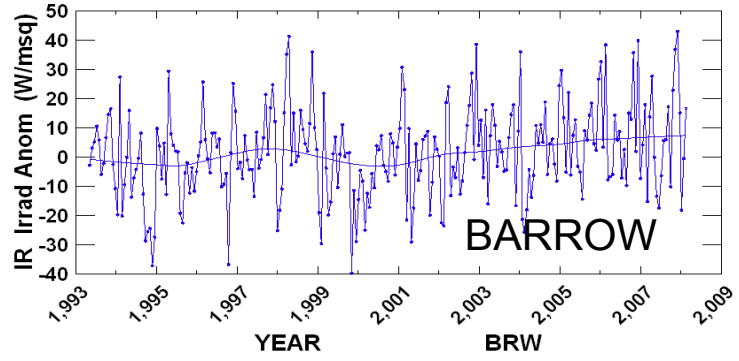
## Surface IR observations and GCM output for grid box containing the site





# 20-day Avg Desasonalized Surface IR Anomalies with Lowess Smoother (0.3)

## ESRL-GMD Radiation Global Baseline Sites



## Linear Trend Detection Times

(required data set duration for detection, B. Weatherhead et al., '98)

Based on:

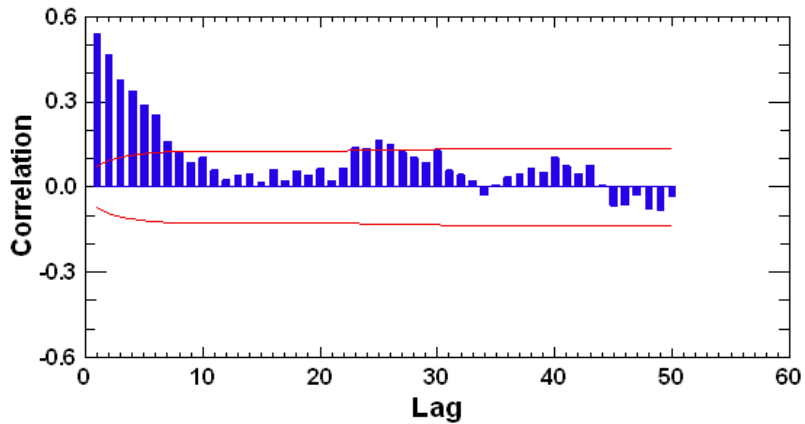
- Estimated variance
- Estimated autocorrelation (AR1)
- Expected trends

For the GMD deseasonalized IR data:

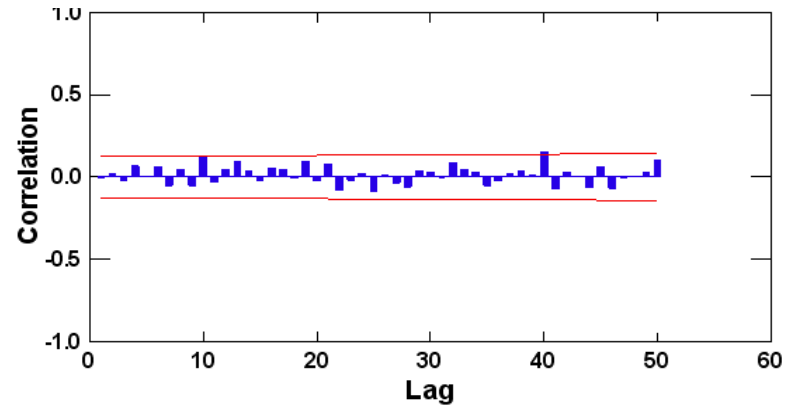
<u>Detectable trend</u>		<u>Uncertainty range in number of required years</u>
0.3 W m <sup>-2</sup> dec <sup>-1</sup>	→	70 to 220 years
2.0 W m <sup>-2</sup> dec <sup>-1</sup>	→	19 to 53 years
3.5 W m <sup>-2</sup> dec <sup>-1</sup>	→	13 to 35 years

Currently have ~15 years of GMD data - It's time to investigate!

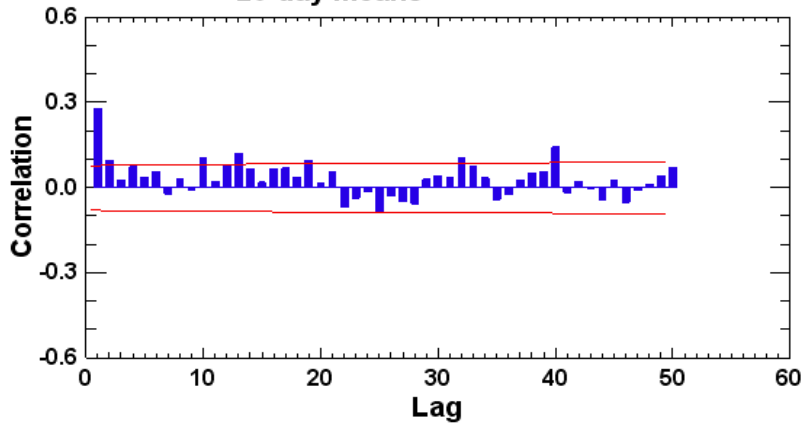
**Autocorrelation Plot - KWAJ**  
20-day means



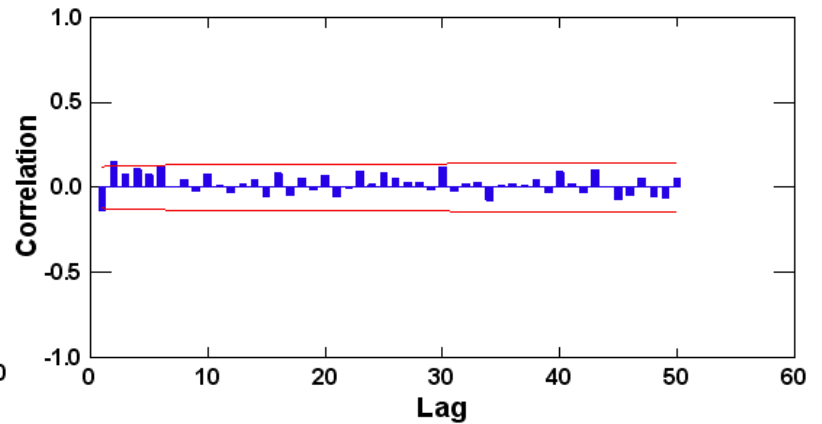
**ARIMA (1,0,0) Residuals**  
Autocorrelation Plot Kwaj



**Autocorrelation Plot - BRW**  
20-day means



**ARIMA (1,0,0) Residuals**  
Autocorrelation Plot BRW



# Estimated Observed Changes in Surface Downward IR

DESEASONIZED  
AR1 Residuals  
Linear trends

Method \ Site	Regress	Mann-Kendall
BRW	5.3	5.2
BLD	3.2	3.4
BER	2.7	2.6
MLO	2.5	1.0
KWA	1.9	1.8
SPO	3.7	2.8
6AVG/SE	3.2/0.5	2.8/0.6
5AVG/SE	3.4/0.6	3.2/0.6

(W m<sup>-2</sup> dec<sup>-1</sup>)

AVG<sub>6</sub> ~ 3.0 (0.6 SE) W m<sup>-2</sup> dec<sup>-1</sup>  
AVG<sub>5</sub> ~ 3.3 (0.6 SE) W m<sup>-2</sup> dec<sup>-1</sup>



Not significant at 95%



Potentially significant at 95%, res uncorrelated, normality tests good to marginal

Avg Regress student's  $t = 2.8$

Avg Mann-Kendall **95% minimum = 0.9 Wm<sup>-2</sup> dec<sup>-1</sup>**

(SPO least sig.)

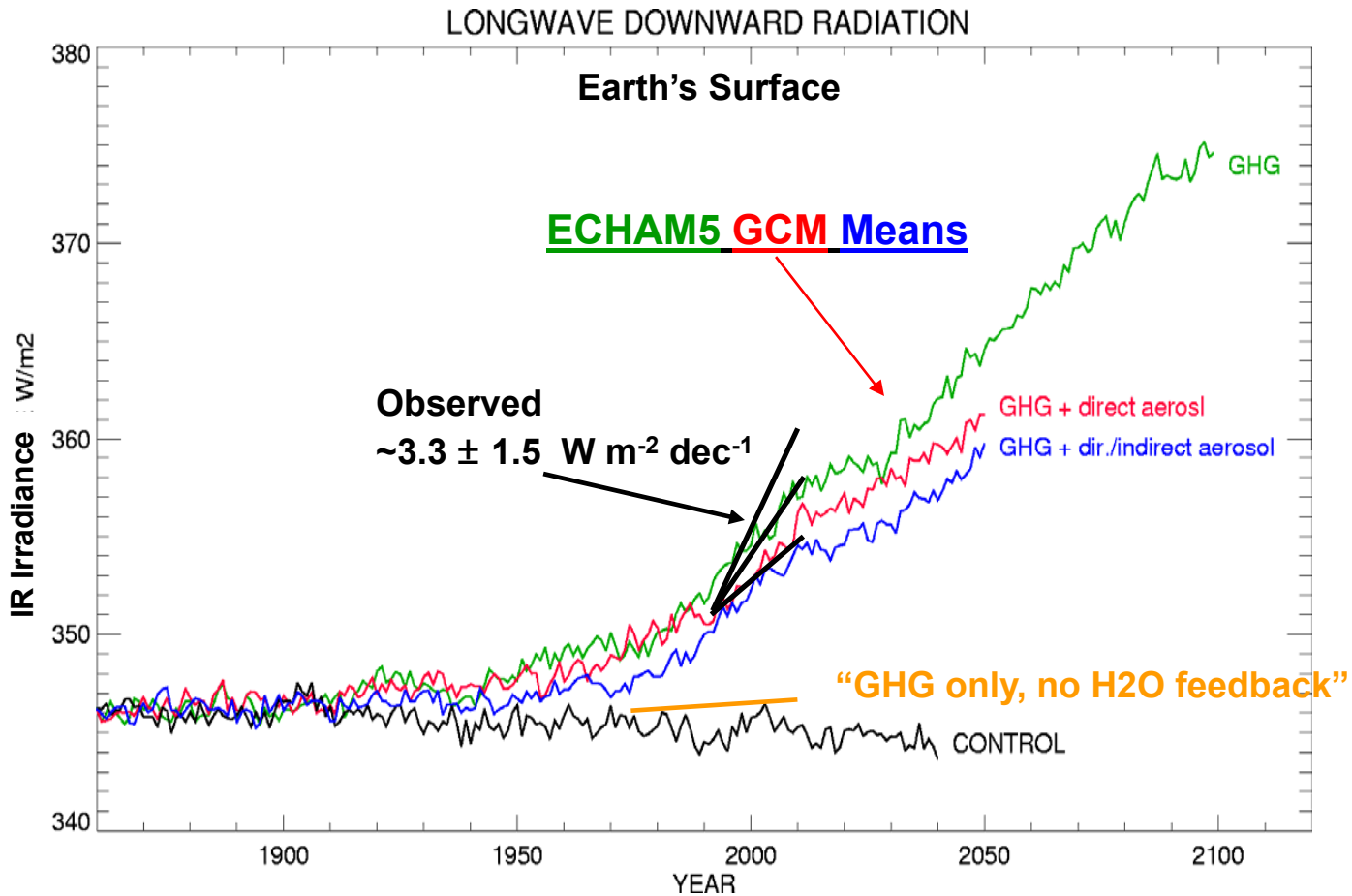
# SURFRAD (CONUS) initial results (Surface IR-down change $W m^{-2} dec^{-1}$ )

Method Site	AR1 res Regress	AR1 res M-K
Ft Peck, Montana	<b>3.3</b>	<b>3.3</b>
Bondville, Illinois	<b>2.5</b>	<b>3.4</b>
Goodwin, Miss.	<b>3.4</b>	<b>2.8</b>

**AVG = 3.1, Boulder (Erie) = 3.3**

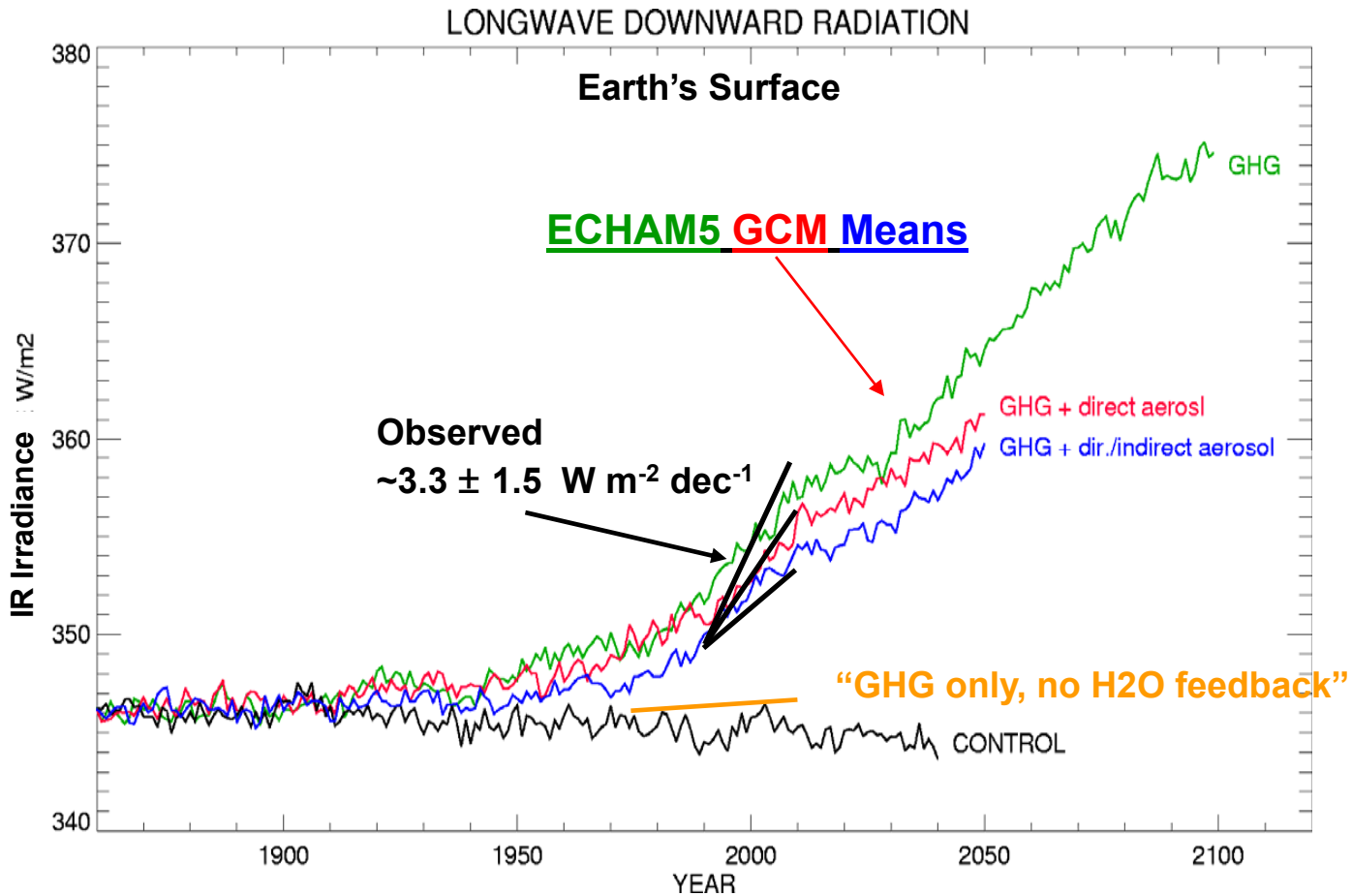
Overall estimate of observed surface downward IR trend  
based on average for five globally remote sites (1993-2008)

**$3.3 \pm \sim 1.5 W m^{-2} dec^{-1}$**



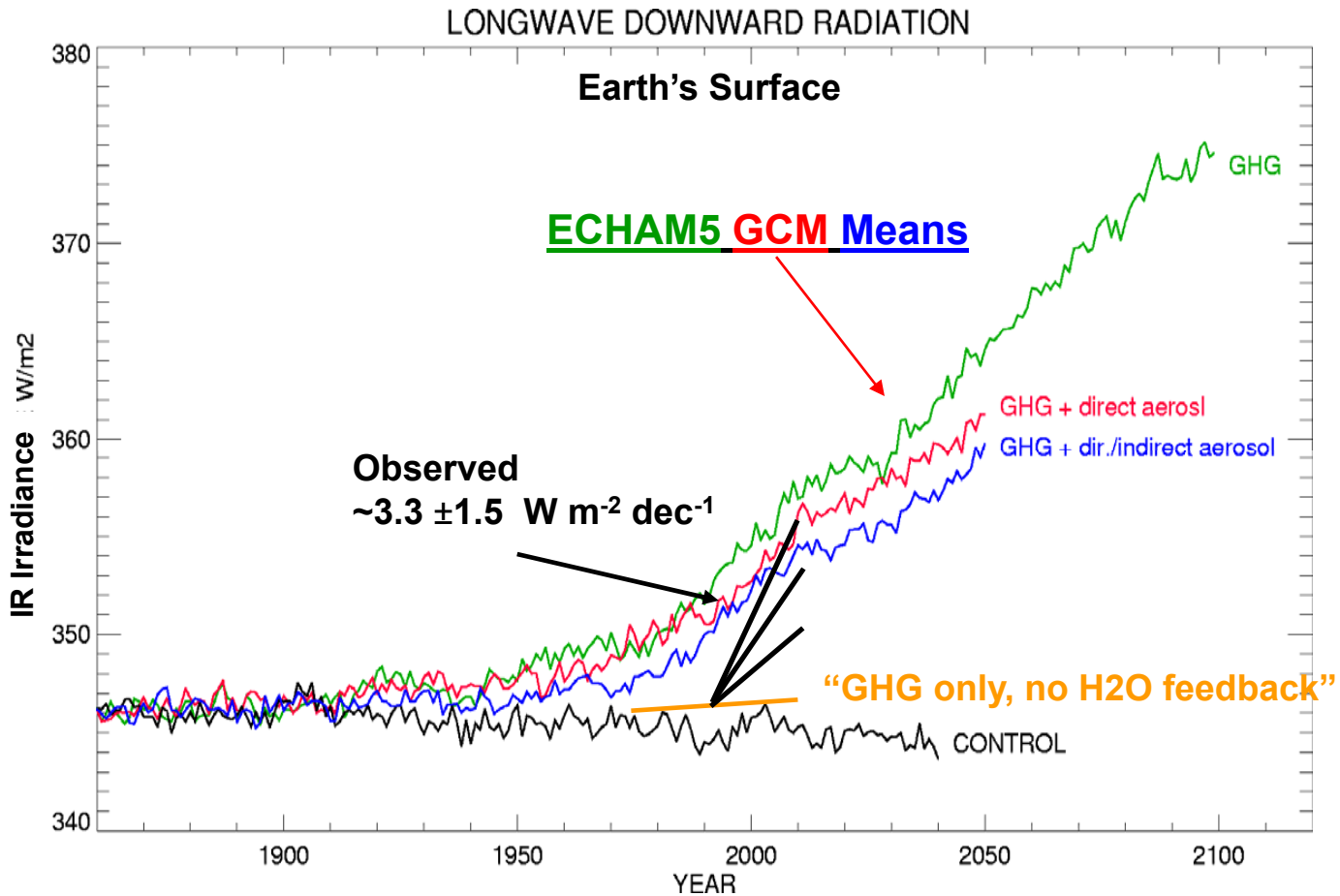
GCM results  
 Provided by  
 Martin Wild / ETHZ

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# Summary

- Using “best estimate” from GMD baseline data, surface IR growing near that predicted by GCMs
- Theoretical statistical estimates of trend detectability are marginally met.
- Maintaining calibration stability and extending the record are crucial
- Mauna Loa is not and should not show as certain a trend as other sites
- The somewhat higher than expected observed growth rates for 1993 – 2008, 3.3 vs 2.5, may be due to Pinatubo cooling recovery and is explicitly consistent with the GFDL fully-forced model run.

## Future plans

- Continue and expand observational effort
- Extend analysis to existing but growing shorter data sets
- More detailed comparisons to fully-forced GCMs in a diagnostic mode
- Adequately determined IR climate could assist in assessing the validity and extent of multiple new and hypothesized feedback mechanisms